Numerical Modeling Of Pressure Source Of Sinabung Volcano Based On GPS Data In 2011-2012 Using Particle Swarm Optimization (PSO)

Ratih Kumalasari¹a, Wahyu Srigutomo², Irwan Meilano² and Hendra Gunawan³

¹Universitas Bhayangkara Jakarta Raya, Jakarta, Indonesia
²Institut Tenologi Bandung, Bandung, Jawa Barat, Indonesia
³Pusat Vulkanologi dan Mitigasi Bencana, Bandung, Jawa Barat, Indonesia

Abstract. Mogi Model with particle swarm optimization (PSO) scheme have been applied to the local GPS data of Sinabung Volcano during 2011 to 2012 to receive subsurface parameters as pressure sources in terms of misfit and inversion model parameter. The size of displacement was inverted by PSO. From the inversion concluded that the position pressure source showing shallow magma pockets at a depth between ±1.3 km volume change around ±0.95×10⁶ m³. It indicates the presence of a huge magma supply and continuous into shallow magma chamber up to the surface of Sinabung Volcano.

Keywords: Pressure source, PSO, Mogi Models, Sinabung

Introduction

Indonesian territory is located at the confluence of several tectonic plates involve the Eurasian Plate, the Indo-Australian Plate and the Pacific Plate. This position causes subduction activity so that many volcanoes are appeared in Indonesia, one of which is Sinabung Volcano. Sinabung Volcano is a strato type volcano that is administratively located in Karo Regency, North Sumatra Province, with a peak of 2,460 m above sea level, with coordinates 3 ° 10 'LU and 98 ° 23.5' BT (PVMBG, 2010).

Before 2010 Sinabung Volcano was a type B volcano, which had not been erupted since 1600s, but still exposed volcanic activities such as the presence of solfatar or fumarole fields. The eruption activities of Sinabung Volcano has just been recorded since August 2010 which make it as type A volcano. Later, continuous monitoring is done to mitigate the danger of the eruptions. Volcanic eruption activity is generally preceded by early symptoms or precursors, such as increased seismic activity, increased temperature of hot springs, changes in composition and strength of gas gusts, as well as the existence of deformations in the body of the volcano. Proper monitoring on the activities of a volcano requires input data from various methods, one of which is a deformation monitoring method by GPS data (Haerani et al., 2012., Kumalasari et al., 2018., Kumalasari et al., 2018.).

Theoretical Background

In this study we used the Mogi model (Point Pressure), Mogi model (1958) assumes that the earth's crust is a semi-elastic medium and the deformation is caused by a point pressure at a
certain depth. If there is a hydrostatic change in the ball, symmetrical deformation will occur. It is explained in the Volcano Deformation that the displacement on the surface caused by changes in hydrostatic pressure in the cavity of the earth's crust with a radius smaller than the depth (\(d \ll a\)) are stated in Equation 3.6 (Dzurisin, 2007)

\[
\begin{bmatrix}
\mu \\ v \\ w
\end{bmatrix} = \alpha \Delta P \left(\frac{1-\nu}{G}\right)
\begin{bmatrix}
x \\ y \\ z
\end{bmatrix}
\]

Where \(d\) is the depth of the pressure source, \(r\) is the radial distance of the pressure source to the monitoring point, \(\nu\) is the poisons ratio, \(G\) is the shear modulus, and \(\Delta P\) is the change in pressure, as graphically illustrated in Figure 7.

\[\text{Figure 1. Mogi Model}\]

Materials and Methods
The data that we used is devide into 2 period:
1. 1st period (1st April 2011 – 22nd July 2011).
2. 2nd period (22nd July 2011 – 4th March 2012).

\[\text{Table 1. Changes of position in each period}\]

<table>
<thead>
<tr>
<th>Period</th>
<th>Station</th>
<th>Coordinate</th>
<th>Displacement (m)</th>
<th>Standard Deviation</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td></td>
<td>x</td>
<td>y</td>
<td>z</td>
</tr>
<tr>
<td>1</td>
<td>LKWR</td>
<td>431687.15000</td>
<td>352794.03000</td>
<td>1496.45190</td>
</tr>
<tr>
<td></td>
<td>SKNL</td>
<td>434757.39000</td>
<td>351034.54000</td>
<td>1442.59190</td>
</tr>
<tr>
<td></td>
<td>GRKI</td>
<td>432671.64000</td>
<td>347790.79000</td>
<td>1230.30610</td>
</tr>
<tr>
<td></td>
<td>SNBGA</td>
<td>440477.08000</td>
<td>347201.59000</td>
<td>1248.57910</td>
</tr>
<tr>
<td>2</td>
<td>LKWR</td>
<td>431687.15000</td>
<td>352794.03000</td>
<td>1496.45190</td>
</tr>
<tr>
<td></td>
<td>SKNL</td>
<td>434757.39000</td>
<td>351034.54000</td>
<td>1442.59190</td>
</tr>
<tr>
<td></td>
<td>GRKI</td>
<td>432671.64000</td>
<td>347790.79000</td>
<td>1230.30610</td>
</tr>
<tr>
<td></td>
<td>SNBGA</td>
<td>440477.08000</td>
<td>347201.59000</td>
<td>1248.57910</td>
</tr>
</tbody>
</table>
We use Particle Swarm Optimization (PSO) to characterize the pressure source parameters of Sinabung volcano. Particle Swarm Optimization (PSO) is an optimization algorithm that mimics the processes in the life survival of a flock of bird and a school of fish developed by James Kennedy and Russell Eberhart in 1995 (Martinez, 2010., Zhou et al., 2018). In PSO, the population is assumed to be a particle with certain size and located at a random location in a multidimensional space. At the initial position of each particle is assumed to have two characteristics, namely position and speed. Each particle moves in a certain space and remembers the best position (pbest) that has ever been passed or found against food sources or the value of an objective function, each particle then conveys its information or best position to other particles and adjusts its position and speed based on information received regarding the best position (gbest). The position and velocity adjustment for each particle is formulated in equation 3.8 and equation 3.9.

\[ V_{ij}(t+1) = wV_{ij}(t) + c_1 r_1(p_{ij} - x_{ij}(t)) + c_2 r_2(p_g - x_{ij}(t)) \]  \[ x_{ij}(t+1) = x_{ij}(t) + v_{ij}(t+1) \]

Which \( V_{ij}(t+1) \) is the speed update that will be used in determining the best position update \( x_{ij}(t+1) \). Then there are constants used in the formulation such as \( w = \text{inertia weight} \), and \( c \) is the velocity coefficient for PSO and \( r \) are random numbers between -1 to 1.

**Results and Discussion**

Results of surface deformation inversion data on Sinabung Volcano using Mogi models with PSO schemes showed good results in misfit and model responses so that it was sufficient to reconstruct the physical field model realistically.

<table>
<thead>
<tr>
<th>Phase</th>
<th>Model</th>
<th>misfit</th>
<th>AIC</th>
<th>A(m)</th>
<th>delta P/G</th>
<th>d(m)</th>
<th>x(m)</th>
<th>y(m)</th>
<th>delta V</th>
</tr>
</thead>
<tbody>
<tr>
<td>1</td>
<td>Mogi dengan PSO</td>
<td>0.0297</td>
<td>10.0593</td>
<td>669.5714</td>
<td>0.0626</td>
<td>27826.6160</td>
<td>450000.00</td>
<td>346001.89</td>
<td>58968083.71</td>
</tr>
<tr>
<td>2</td>
<td>Mogi dengan PSO</td>
<td>0.0564</td>
<td>10.8718</td>
<td>236.1255</td>
<td>0.0230</td>
<td>1300.0000</td>
<td>428500.00</td>
<td>348148.75</td>
<td>952446.95</td>
</tr>
</tbody>
</table>

Table 2. Result of Parameters Model Using PSO
The inversion results in period 1 showed inversion result at depth of around 27.8 km with change in the volume around $3.61 \times 10^6 \text{m}^3$. This result is possible that the source of the inversion pressure is not a local pressure source from Sinabung Volcano but another pressure source which can be a tectonic or partial melting source which requires further study. In period 1 Sinabung Volcano was classified in a normal active period where no eruptions occurred and the status of the Sinabung Volcano was downgraded after eruptions his first eruption from 27 August 2010 to 3 September 2010 (Gunawan, 2017), that condition also reinforces that the deformation of the surface of Sinabung Volcano in period 1 is not from the local source.
The inversion results in period 2 show the smallest misfit obtained from the inversion in the Mogi model with the PSO scheme. The source of the inversion pressure is at a depth of ±1.3 km with volume change around ±0.95×10^6 m^3 in the northeast sector of Sinabung Volcano. With a fairly shallow in depth of the pressure source in period 2, it can concluded that the source is magma pocket of Sinabung Volcano.

The inversion results are selected based on data misfit in accordance with the results of tectonic seismic hypocenter relocation by Indrastuti (2017) shown in Figure 20. Based on the inversion of the pressure source elaborated by the relocation of tectonic volcanic activity, it is known that tectonic volcano is associated with magma pocket of Sinabung Volcano, from the model also known that there are magma supply from deep magma pocket to moderate magma pocket and shallow magma pocket which marked by seismic hypocenters while at the same time strengthening that the inversion results are in accordance with the conditions of Sinabung Volcano.

![Figure 4. Hypocenter relocation of volcano tektonik (Indrastuti, 2017)](image)

Conclusions

The model resulting from the inversion process is sufficient to reconstruct the physical field. From the inversion concluded that the position pressure source showing shallow magma pockets at a depth between ±1.3 km volume change around ±0.95×10^6 m^3. It indicates the presence of a huge magma supply and continuous into shallow magma chamber up to the surface of Sinabung Volcano.
References


